Volatile rich Fe-oxide melts: base and precious metals enrichment trends. Is there an IOCG connection?

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Background



H.R. Dick Naslund Binghamton University

Nelsonites and Low Ti, FeO deposits in the Adirondacks, New York State

El Laco, Chile and the Missouri, USA FeO-REE deposits, with Fernando Enrique and Jan Nystrom

6 years Alan Clarks IOCG group, Queens University, Canada

Jorge Benivides- Manto Verde, Chile Hyong Chen- Marcona, Mina Justa, Peru Greg Lester- Pampa De Pongo, Peru Proprietary deposits in Mexico, Canada and 100+ million pound REO, IOCG-carbonatite hybrid in Mongolia

Observations

- Magmatic-hydrothermal systems
- Usual variations of the typical IOCG alteration/ alteration/mineralogy themes
- Primary early FeO stage volatile rich AAFEs (anions or anion forming elements) e.g. OH, P, S, F
- External or non-primary source of S, and by extension Cu and Au
- REE and U might be primary
- Geochemical and textural evidence of primary FeO magmas in some cases

Experimental approach If liquid immiscibility between volatile-rich FeO and SiO₂ dominated melts might be involved in the genesis of IOAs or IOCGs, what should be determined and considered?

- Pressure, temperature and chemical stability fields of two melt systems.
- Determine the partitioning of major elements, REEs, metals, and other trace elements, and stable isotopes.
 The melts should be H₂O-rich and characterize the effects of P, S, Cl or F on phase stability and element enrichment
 Variable fO₂, QFM, NNO, MH

>200 experiments

T-X-7 base mixtures (~basaltic and esite-granite comp.) + H_2O +P,S, F, or S

Part.- base-mixture + 10 wt% H_2O + 6.67 x 10⁻⁴ mol/g P, S, F, or Cl + trace elements (peraluminous, An_{50})



Experimental configuration

2mm Pt inner capsule

> External buffer + H₂O (QFM, MH, NNO)

Experimental charge: Si-Al-K-Fe-O + H_2O+P , S, F or Cl, + trace elements



Stellite bomb

Rotating, rapid-quench IHPV

Argon pressure media Argon density at 200 MPa ~ density water at 1 atm.



Quench orientation



blast shield ⁻

Run orientation

Experimental product for analysis: phase assemblage (microscopy, B.S.E), major and trace element (W.D.S., ICP-ms) and isotope (IR-ms) partitioning



trans. light



refl. light

B.S.E. images of conjugate immiscible melts









Experimental results for $H_2O + P$



L - liquid

M - magnetite

Sil - silica mineral

Experimental results for $H_2O + S$



L - liquid

M - magnetite

Sil - silica mineral

Experimental results for $H_2O + F$



L - liquid

M - magnetite

Sil - silica mineral

Element Partitioning as a Function of Polymerization



Effect of H₂O, P, S, and F on miscibility gap



<u>Summary of results and interpretation-P-T-X stability</u>
1. FeO_x-SiO₂ volatile rich immiscible melts are stable over a broad melt composition range, at geologically reasonable temperatures 1000-1200 C° (800 C° sub-liquidus). 60-80 wt% FeO FeO-rich melt, andesitic-granitic silicate melt

2. H₂O, P, S, F extend T lower, expand composition field, Cl-limited

Mechanisms

Crystal fractionation, magma mixing, partial melt, assimilation

Geological Settings

High H_2O Arc and back-arc are very permissive, especially with deep faults as conduits.

Shallow crustal magma chambers (late LLD, volatile saturation near and sub-liquidus, also magma mixing)

Mid cratonic- Volatile source? Low volatile IO?, IOA

Pressure and two-melt stability in the supra-subduction zone. "magma factory"



Partitioning of HFSE, transition metals and REE between immiscible silicate melts with H₂O, P, S, F, and CI General applications



Three suites of trace elements

<u>Alk, HFSE</u>: Rb, Cs, Nb, Zr, Cs, Hf, Th, U

<u>REE</u>: La, Nd, Ce, Sm, Eu, Dy, Er, Yb

<u>VI-XII elements</u>: Ti, V, Cr, Mn, Co, Ni, Cu, Zn, Mo, Ag, Au

Experimental conditions

- Temperature: isothermal- 1200 °C
- Pressure: isobaric- 200 MPa
- Oxygen fugacity: fixed at quartz-fayalitemagnetite, nickel-nickel oxide and magnetitehematite buffers
- Molar equivalent quantities of P, S, F, and Cl.

Analysis and Results

Major element: Electron microprobe, WDS

Trace element: Thermo Scientific XSERIES 2 Quadrupole ICPMS instrument, coupled with a NewWave 213 nmlaser ablation system. The laser beam was fired at 5 Hz repetition rate with the laser energy in the range of 0.36–0.51 mJ per pulse. A spot size of 30 μ m.

Nernst partition coefficients D (L^m/L^f) for HFSE and Alk. in conjugate immiscible melts, note log scale



Nernst partition coefficients D (L^m/L^f) for REE in conjugate immiscible melts



Nernst partition coefficients D (L^m/L^f) for VI-XII elements in conjugate immiscible melts



Nernst partition coefficients D (L^m/L^f) for VI-XII elements in conjugate immiscible melts



What about gold?

Summary

 Volatile –rich magma un-mixing is an efficient mechanism for generating IOs, IOAs, maybe IOCGs, in part.

 P- extreme REE and U enrichment • S- three liquids at high fO₂ • F, P, S and Cl with or without H₂O act to rge SiO₂ T-X stability field thus enriching residual melts with desirables. These systems are saturated or near saturation with regard to many desirable elements, FeO redox role, thus fertile loci for external fluids. Fe-oxide deposit -El Laco, Chile, el. 5200 m, (17,500 ft)

Exploration, further considerations and tidbits

- Matches what we already know, and nothing says "Hi" like alteration and FeO targeted geophysics!
- Emulsions
- Melt-fluid partitioning- what would we expect? further work
 Nelsonitic-carbonatitic-REE connection, very
 - important, F and P

Kiruna, Sweden

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Queens University Facility for Isotope Research Further interest: Lester et al.; Chemical Geology, 2013, Contributions to Mineralogy and Petrology, 2013, Geochim Cosmochim Acta, 2013